

Space weather impact in the D-region ionosphere: observations and modelling

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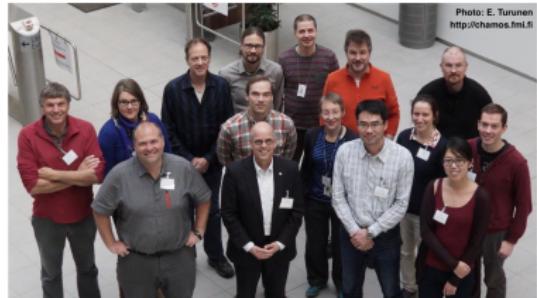
2 May 2024

CHAMOS (<http://chamos.fmi.fi>)

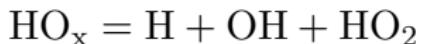
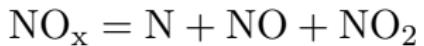
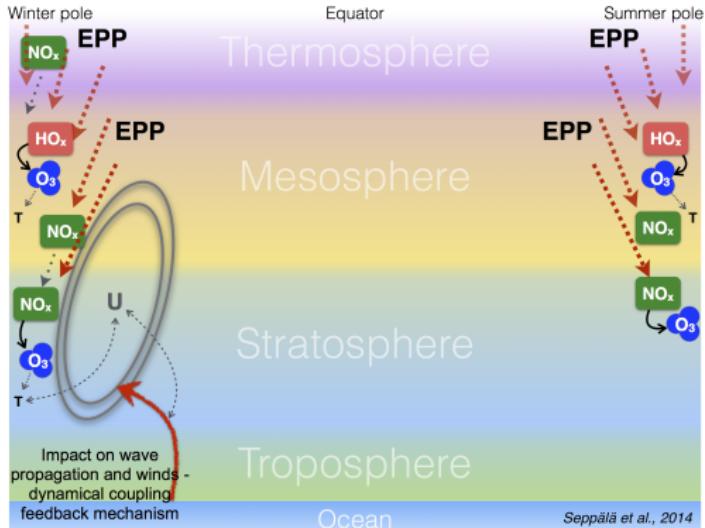


Main page

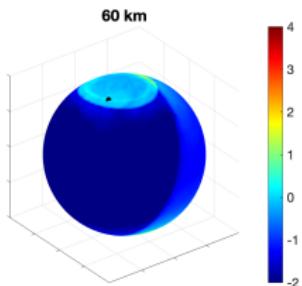
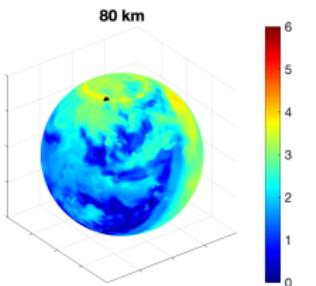
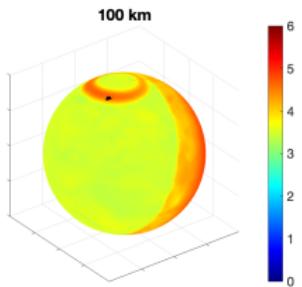
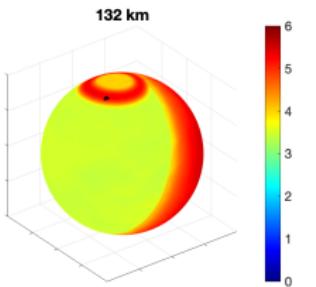
Last modified: 08-Jan-2019



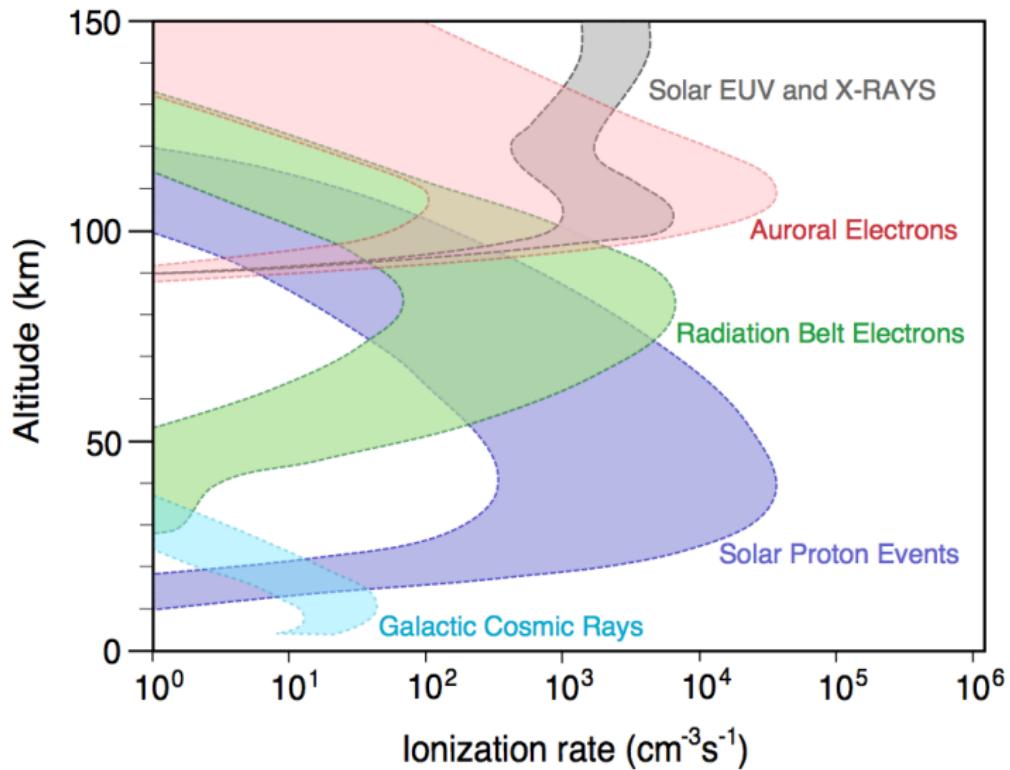
Chemical Aeronomy in the Mesosphere and Ozone in the Stratosphere (CHAMOS) is an on-going science collaboration focusing on energetic particle precipitation and related atmospheric effects. Originally, it was started by scientist from Finnish Meteorological Institute (FMI) and Sodankylä Geophysical Observatory (SGO) back in



WACCM-D model video

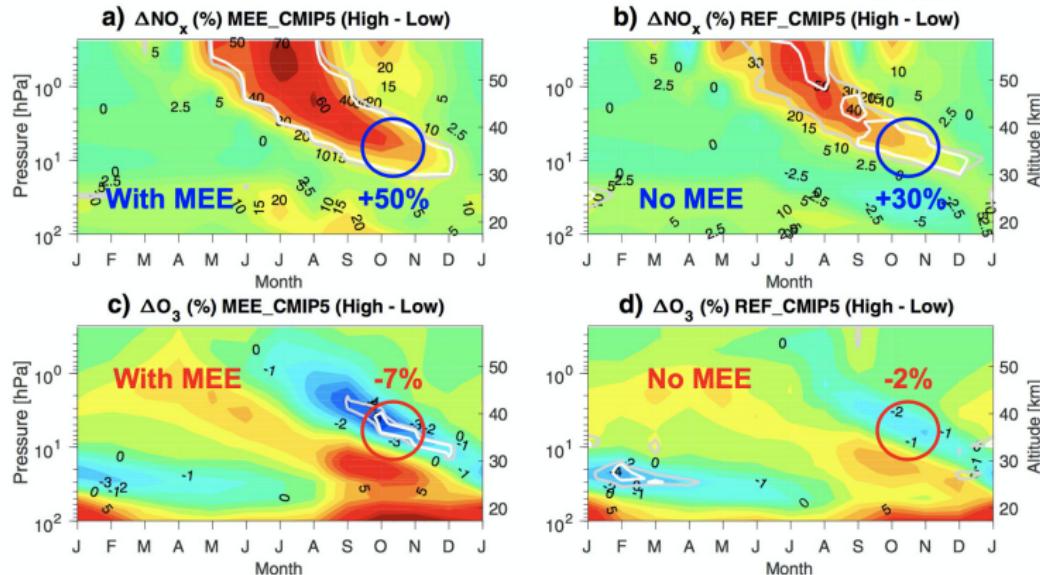


Ionisation sources



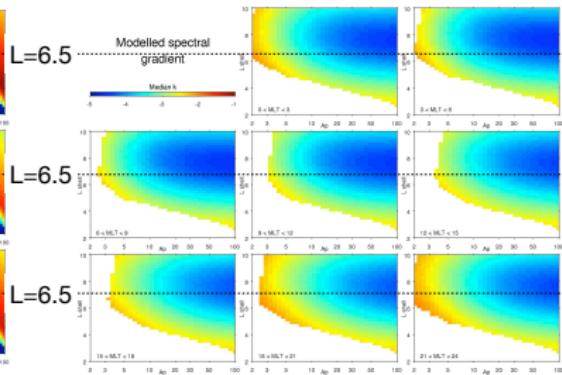
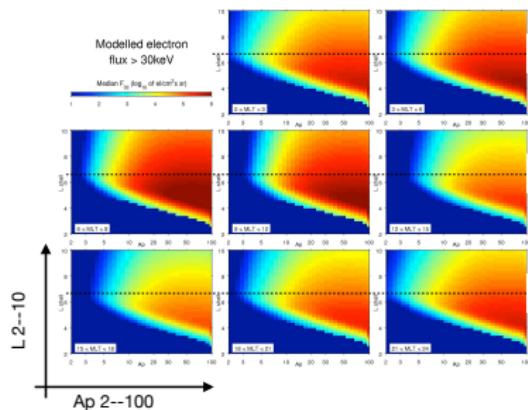
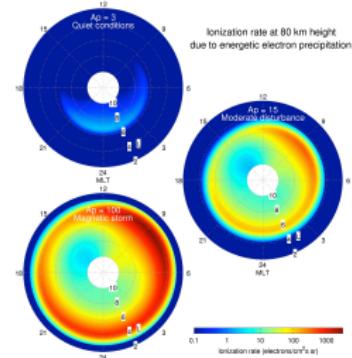
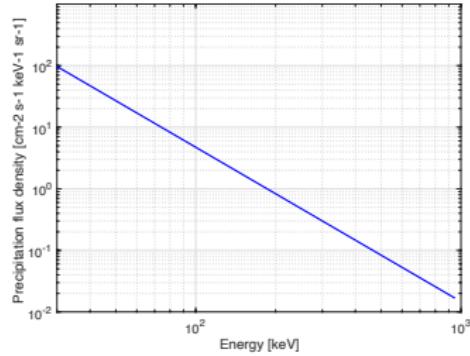
CMIP with (and without) the medium energy electrons (E=30...1000 keV)

Stratosphere 60°S - 90°S



From: Andersson et al., JGR (Atmos.), <http://doi.org/10.1002/2017JD027605>, 2018

APEEP precipitation model, van de Kamp et al., 2016



Tyssøy et al., 2021: HEPPA III intercomparison experiment on electron precipitation impacts



Journal of Geophysical Research: Space Physics

10.1029/2021JA029128

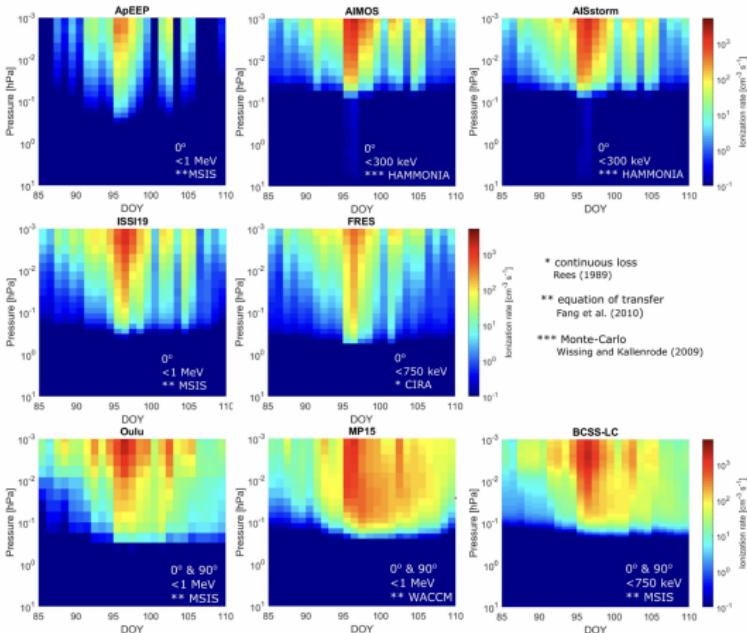


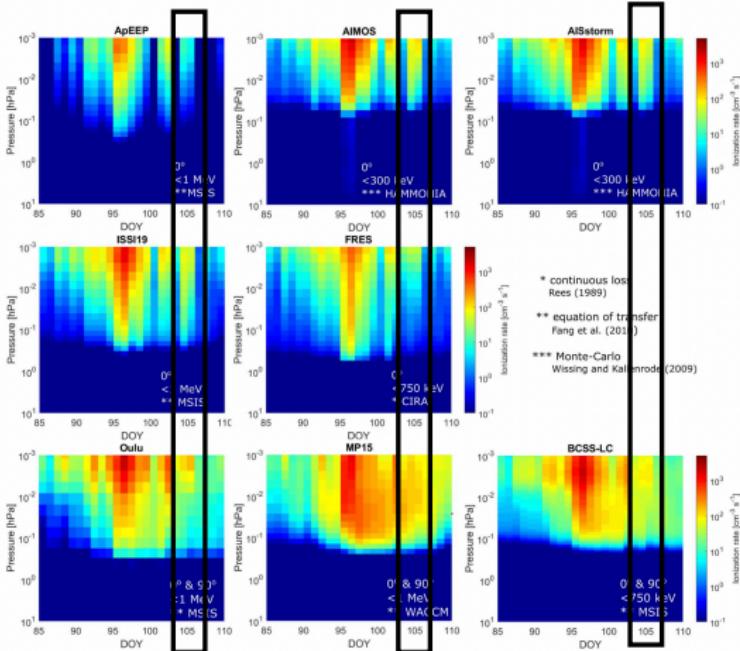
Figure 4. Latitude corrected hemispheric mean poleward of 45°S for the eight ionization rate estimates. The legends list the detector(s), upper energy limit, background atmosphere and ionization rate method applied.

EISCAT data available for 15th April 2010 (doy 105)!



Journal of Geophysical Research: Space Physics

10.1029/2021JA029128



¹<https://doi.org/10.1029/2021JA029128>

EPP inversion:

$$\phi(E) \rightarrow Q(h) \rightarrow N_e(h)$$

① $\phi(E) \rightarrow Q(h)$

- ▶ Calculate $Q(h) = \int A(h, E) \phi(E) dE$
- ▶ Matrix $A(h, E)$ contains the atmospheric response¹

② $Q(h) \rightarrow \text{SIC} \rightarrow N_e(h)$

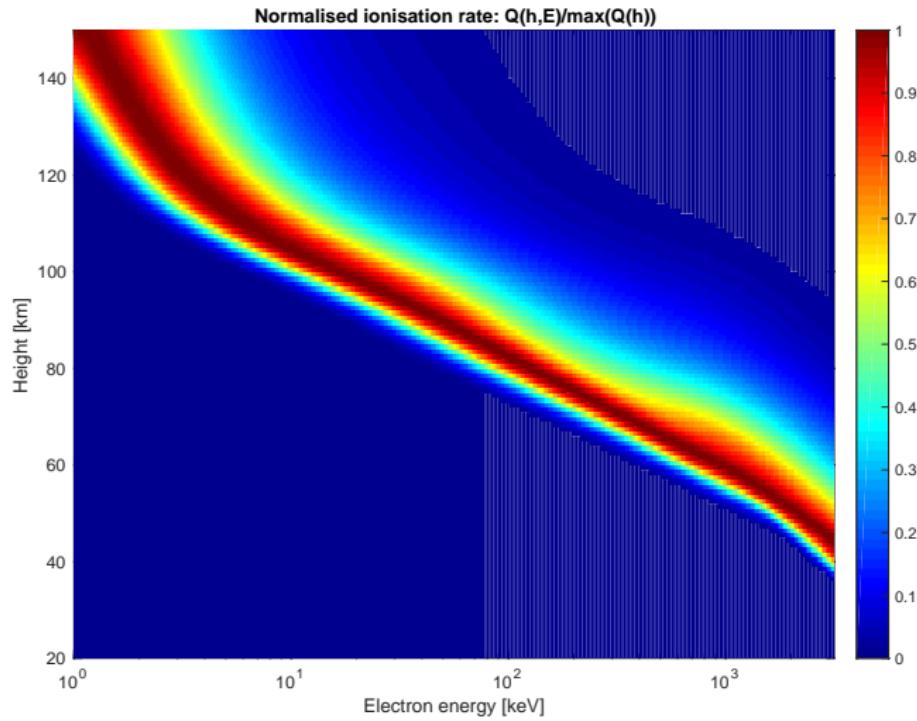
- ▶ Pre-calculate N_e at a sufficiently fine grid of Q and interpolate $Q \rightarrow N_e$

③ Inversion

- ▶ Proposal $\phi(E)$ is an arbitrary piece-wise log-log (here: 7 nodes)
- ▶ Minimise in MCMC chain ($N=100\ 000$):
$$\chi^2 = \sum_h ((\log(N_e(h)) - \log(N_{e:\text{data}}(h))) / W)^2$$

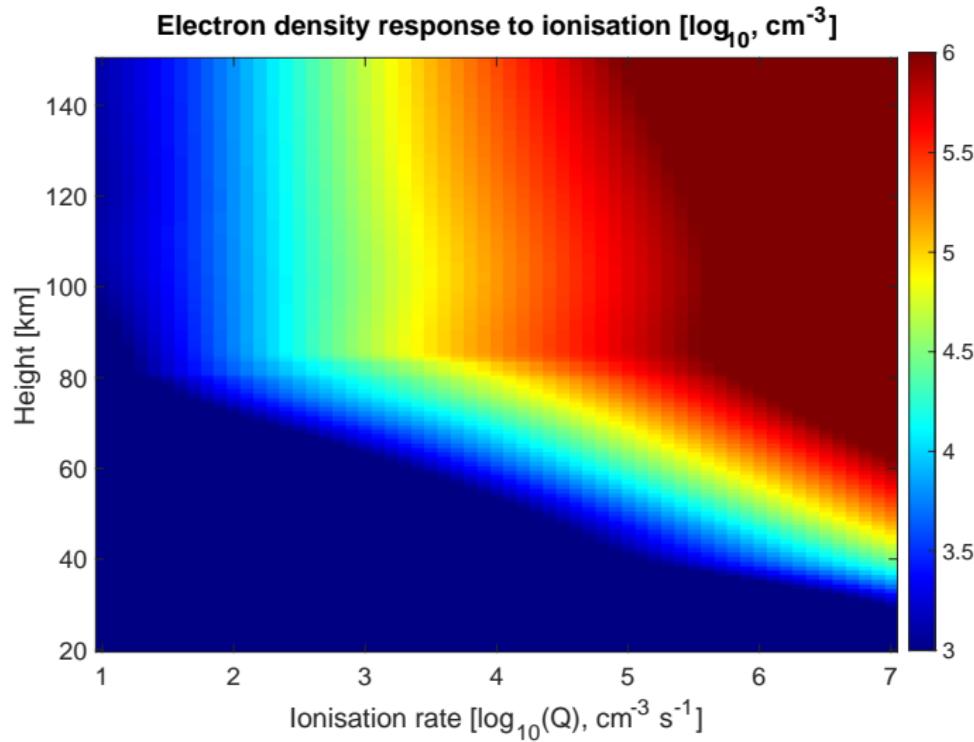
¹Fang, X., C. E. Randall, D. Lummerzheim, W. Wang, G. Lu, S. C. Solomon, and R. A. Frahm (2010), Parameterization of monoenergetic electron impact ionization, Geophys. Res. Lett., 37, L22106, doi:10.1029/2010GL045406.

EPP: energy vs. ionisation

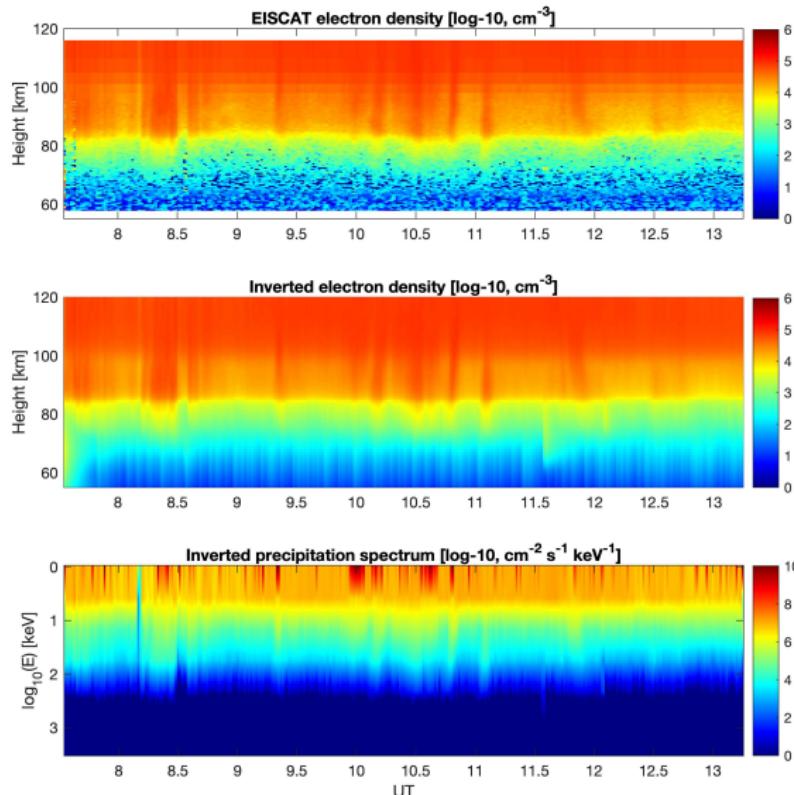


¹Fang et al., Parameterization of monoenergetic electron impact, GRL, 2010

Ionisation impact on electron density (SIC model)

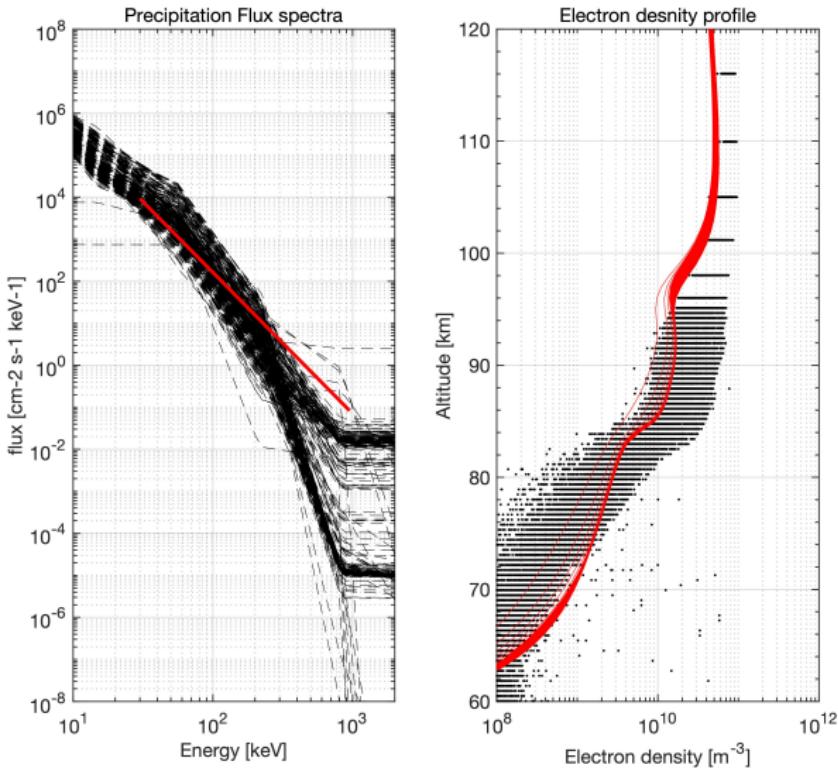


Inversion analysis of EISCAT VHF, 15 April 2010



Comparison: EISCAT vs. APEEP

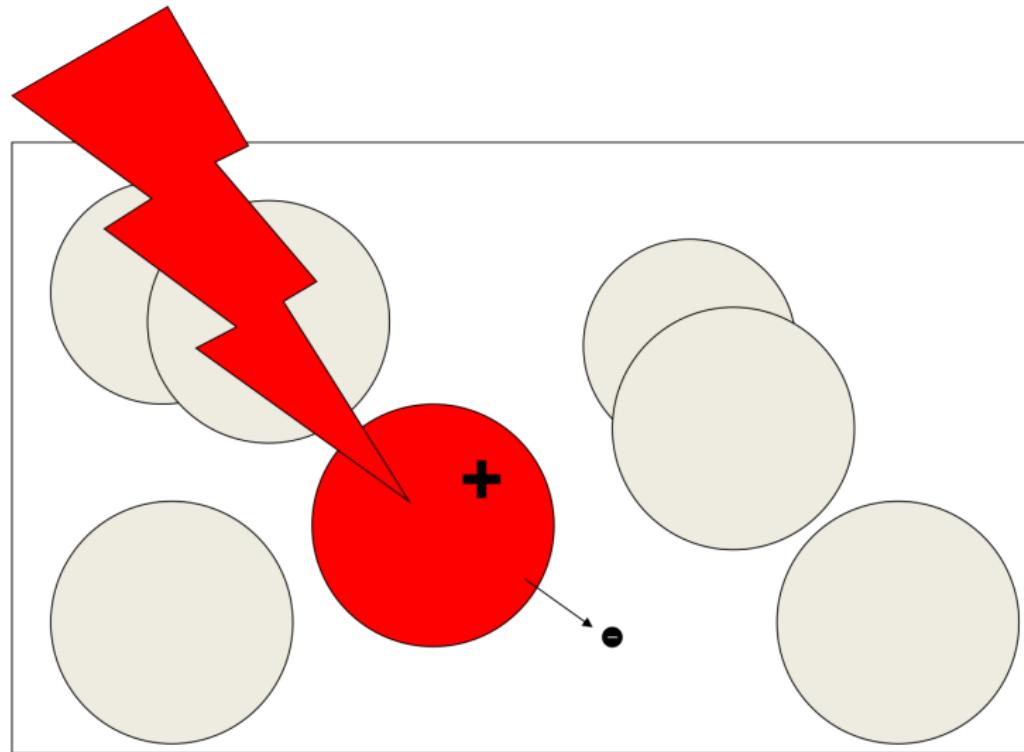
APEEP vs EISCAT 15-April-2010



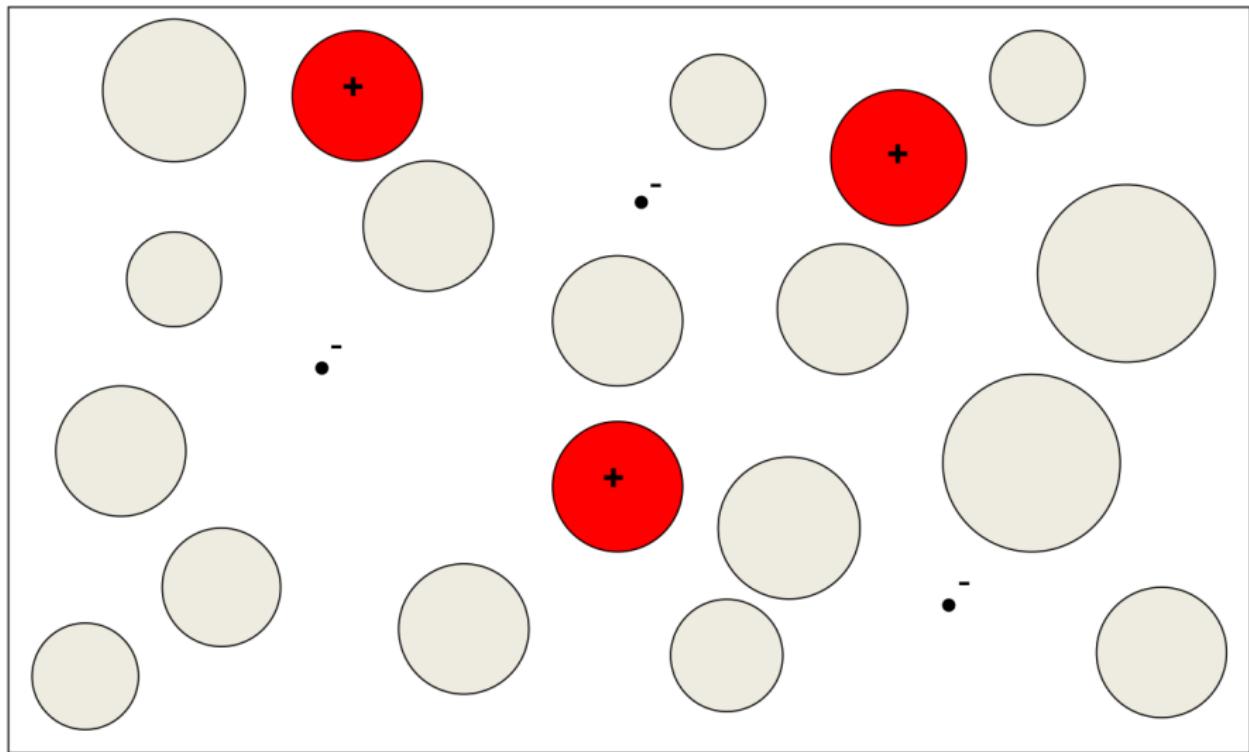
Summary of part 1/3

- APEEP seems to be in *the right ballpark* with the EISCAT data
- Slight overestimation at the relativistic fluxes is likely due to "stiffness" of the spectral shape
- This case study does not seem to provide an easy solution for the NOx dilemma, i.e., there's *no drastic underestimation* of the ionisation.
- All EISCAT D-region data available for similar comparisons to validate & improve the model.

Part 2/3: Modelling of consequences of ionisation



Ionisation → plasma



Building a chemistry model

Chemical reactions introduce a set of differential continuity equations:
The most simple: just 1 positive ion type N_{ip} :

$$\frac{dN_e}{dt} = Q - k_1 N_e N_{ip} = Q - k_1 N_e^2 \quad (1)$$

For one type of positive N_{ip} and one type of negative N_{in} ions:

$$\frac{dN_e}{dt} = Q - k_1 N_e N_{ip} - k_2 N_e + k_3 N_{in} \quad (2)$$

And so forth ...

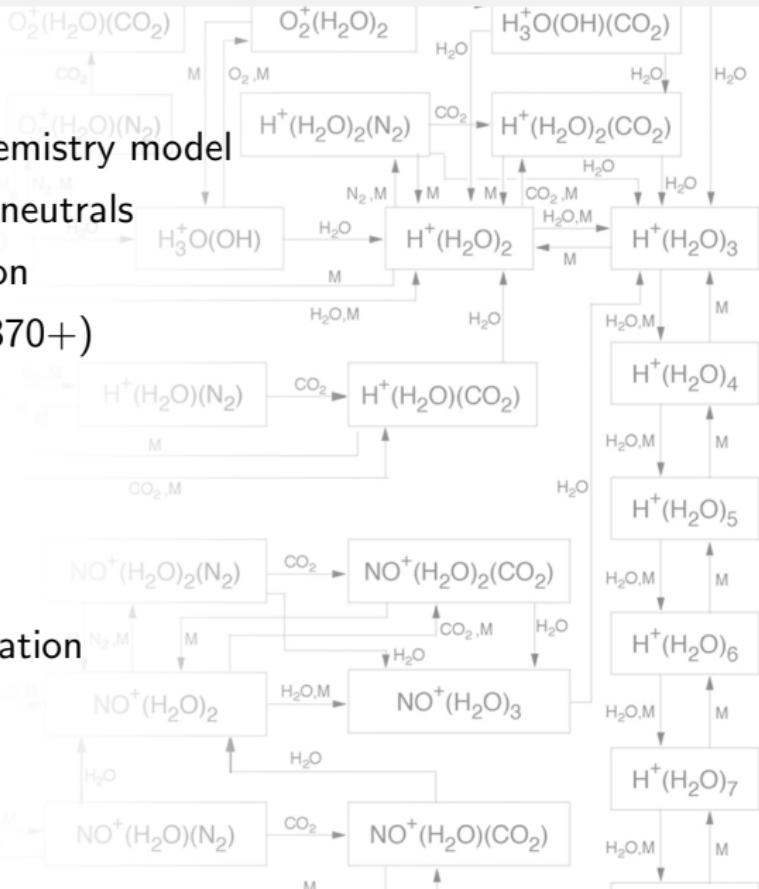
Sodankylä Ion (and neutral) Chemistry model (SIC)

Detailed 1-D time dependent chemistry model

- 63 ions (27 negative) & 13 neutrals
- 20-150 km in 1 km resolution
- several hundred reactions (370+)
- vertical transport

Input

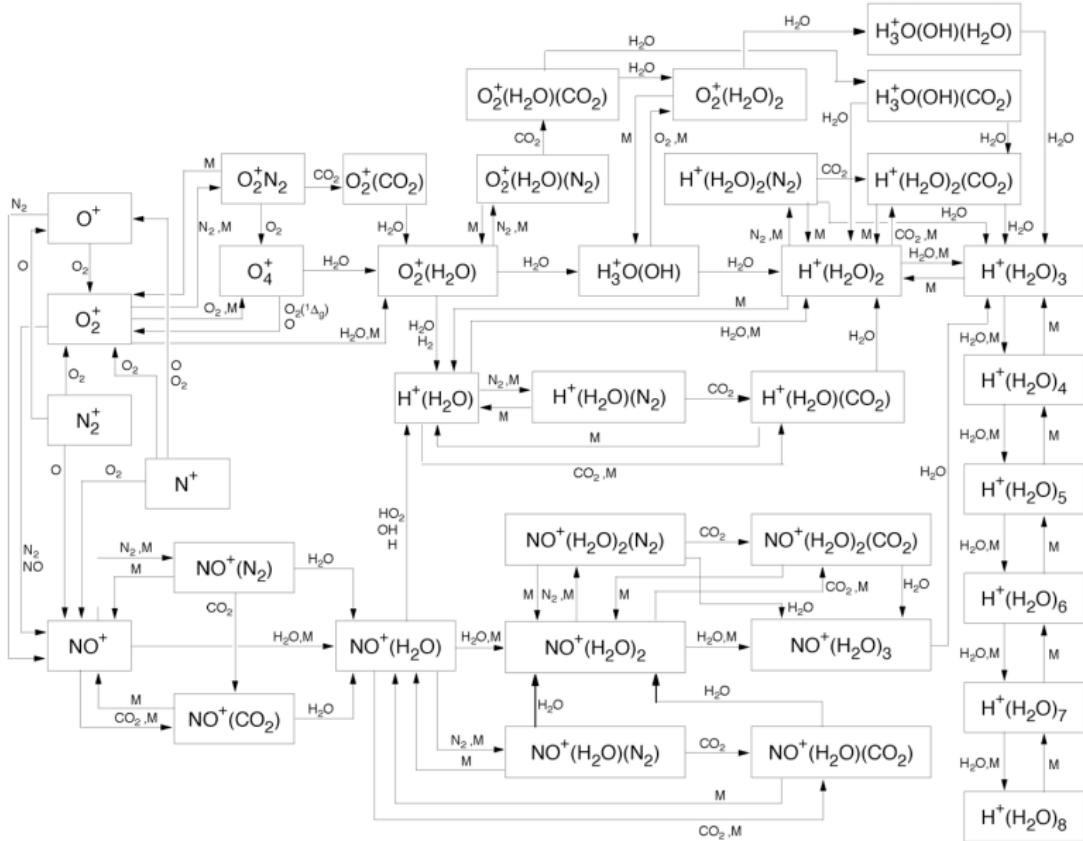
- MSIS
- solar EM flux
- proton and electron precipitation
- cosmic rays



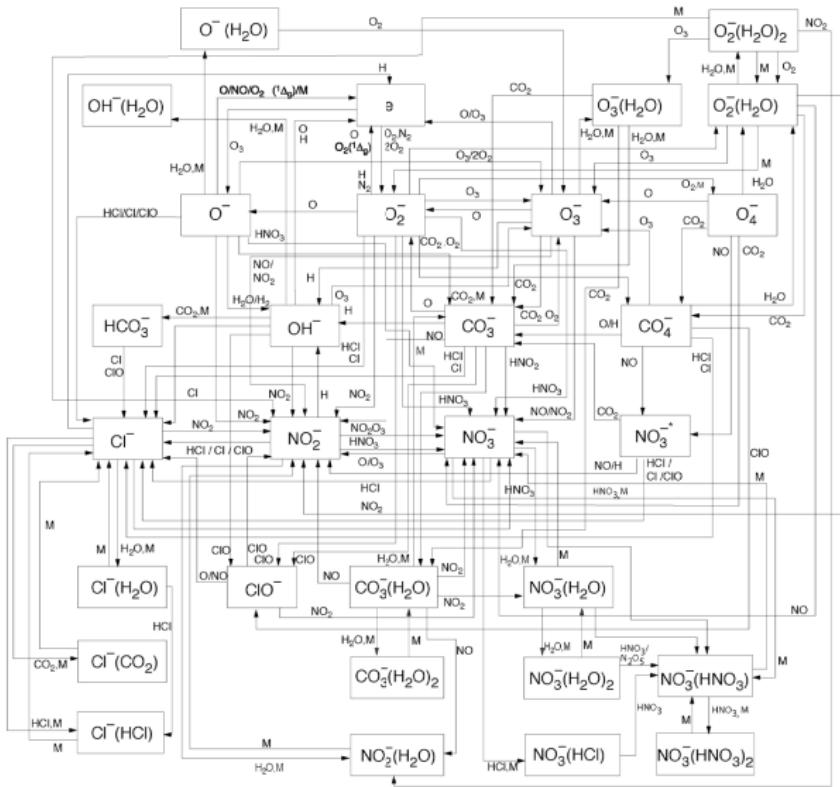
SIC species

POSITIVE IONS		NEGATIVE IONS	NEUTRALS
1	O ⁺	1	O ⁻
2	O ₂ ⁺	2	O ₂ ⁻
3	O ₄ ⁺	3	O ₃ ⁻
4	N ⁺	4	O ₄ ⁻
5	N ₂ ⁺	5	OH ⁻
6	NO ⁺	6	CO ₃ ⁻
7	NO ^{+(N₂)}	7	CO ₄ ⁻
8	NO ^{+(CO₂)}	8	NO ₂ ⁻
9	NO ^{+(H₂O)}	9	NO ₃ ⁻
10	NO ^{+(H₂O)₂}	10	NO ₃ ^{*-}
11	NO ^{+(H₂O)₃}	11	HCO ₃ ⁻
12	NO ^{+(H₂O)(N₂)}	12	O ₂ ⁻ (H ₂ O)
13	NO ^{+(H₂O)(CO₂)}	13	Cl ⁻
14	NO ^{+(H₂O)₂(N₂)}	14	OH ⁻ (H ₂ O)
15	NO ^{+(H₂O)₂(CO₂)}	15	O ₃ ⁻ (H ₂ O)
16	O ₂ ^{+(H₂O)}	16	NO ₃ ⁻ (H ₂ O)
17	H ₃ O ^{+(OH)}	17	CO ₃ ⁻ (H ₂ O)
18	H ^{+(H₂O)}	18	Cl ⁻ (H ₂ O)
19	H ^{+(H₂O)₂}	19	ClO ⁻
20	H ^{+(H₂O)₃}		20 H ₂
21	H ^{+(H₂O)₄}		21 HCl
22	H ^{+(H₂O)₅}		22 HNO ₃
23	H ^{+(H₂O)₆}		23 Cl
24	H ^{+(H₂O)₇}		24 ClO
25	H ^{+(H₂O)₈}		25 CH ₄
26	O ₂ ^{+(N₂)}		26 CH ₃
27	H ^{+(H₂O)₂(CO₂)}		27 O(¹ D)
28	H ^{+(H₂O)₂N₂}		28 N ₂ O
29	H ^{+(H₂O)CO₂}		29 N(² D)
30	H ^{+(H₂O)N₂}		30 H ₂ O ₂
31	O ₂ ^{+(CO₂)}		31 CO
32	O ₂ ^{+(H₂O)N₂}		32 CH ₂ O
33	O ₂ ^{+(H₂O)CO₂}		
34	O ₂ ^{+(H₂O)₂}		
35	H ₃ O ^{+(OH)H₂O}		
36	H ₃ O ^{+(OH)CO₂}		

Positive ion chemistry scheme



Negative ion chemistry scheme



Solving an arbitrary chemistry scheme

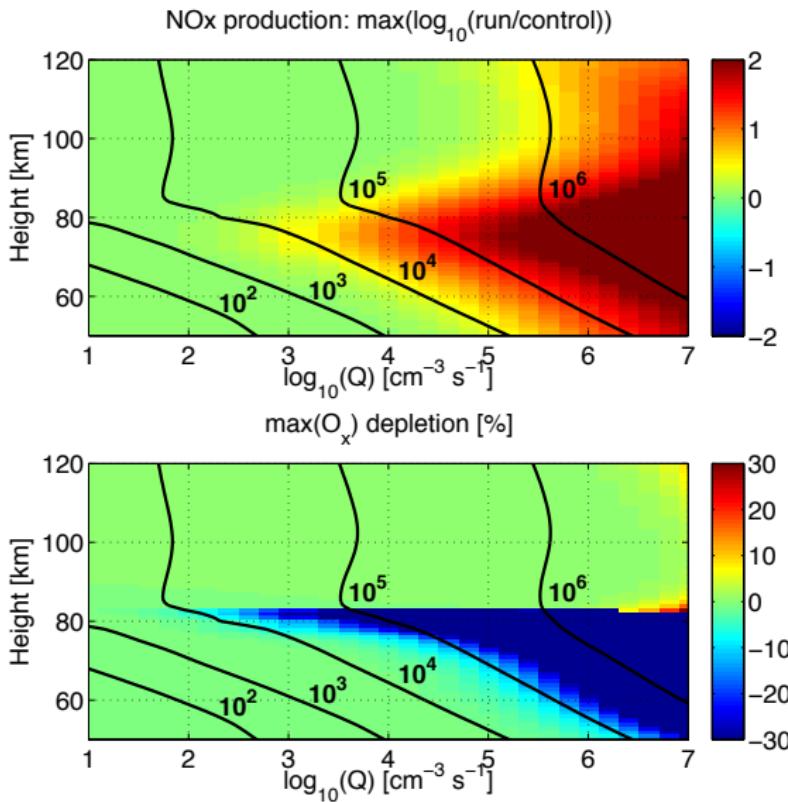
$$\frac{\partial \bar{N}}{\partial t} = \mathbf{B}\bar{N} + \bar{Q} = \begin{pmatrix} -\Lambda_1 & +\Pi_{12} & \dots & +\Pi_{1n} \\ +\Pi_{21} & -\Lambda_2 & & \\ \vdots & & \ddots & \\ +\Pi_{n1} & & & -\Lambda_n \end{pmatrix} \begin{pmatrix} n_1 \\ n_2 \\ n_3 \\ \vdots \\ n_n \end{pmatrix} + \begin{pmatrix} q_1 \\ q_2 \\ q_3 \\ \vdots \\ q_n \end{pmatrix}. \quad (4.9)$$

¹Courtesy of Verronen, 2006

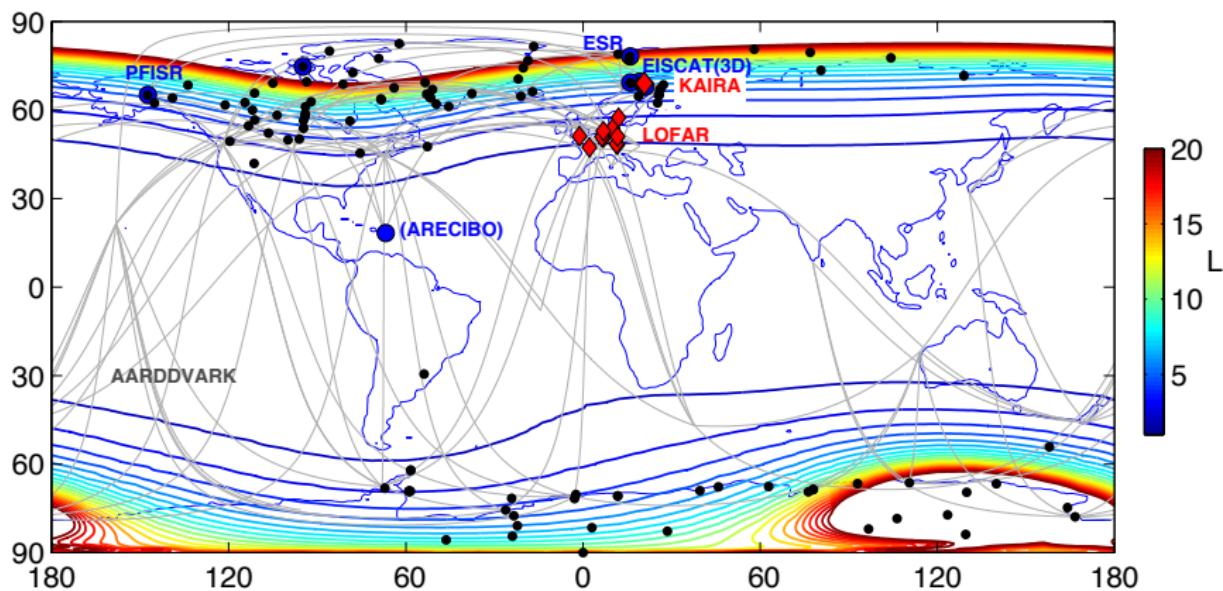
From EPP to NO_x and HO_x

- EPP produces O₂⁺ (also: N₂⁺ + O₂ → O₂⁺ + N₂)
- NO_x is produced e.g., via: O₂⁺ + N₂ → NO⁺ + NO
- HO_x: O₂⁺ turns into water clusters O₂⁺(H₂O)
- Recombination with electrons turn O₂⁺(H₂O) into OH
- Recombination with negative ions turn O₂⁺(H₂O) into HNO₃, which is photodissociated into OH

The end of part 2/3: Sensitivity to ionisation (30 minutes)



Part 3/3: Radio wave methods for observing the ionisation (=electron density)?



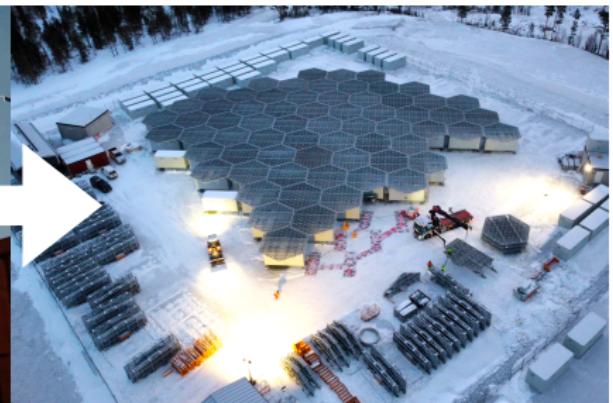
Incoherent scatter radars: paradigm change

Current EISCAT



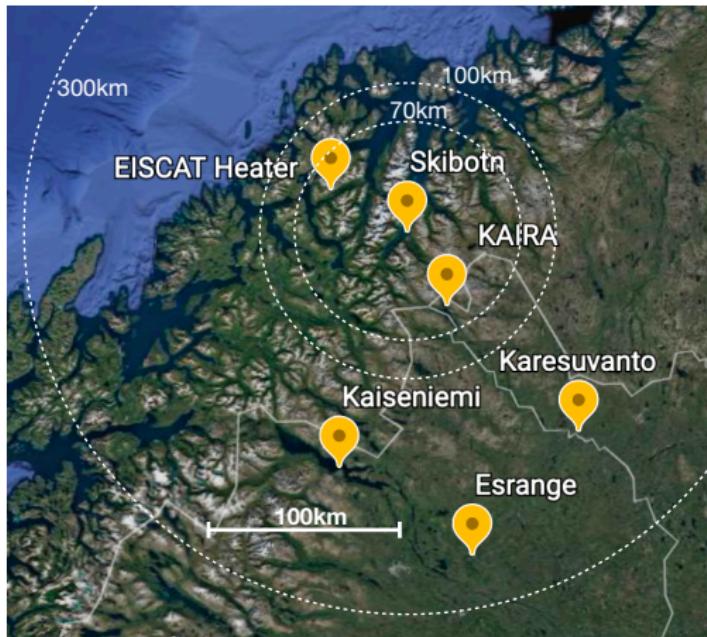
UHF-radar at Tromsø, Norway

EISCAT_3D

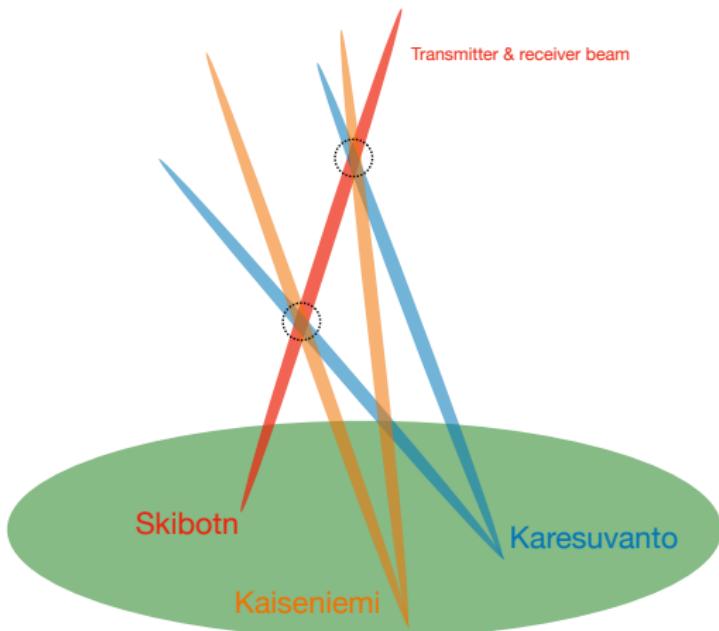


EISCAT_3D transmitter site
Skibotn, Norway

EISCAT_3D geography



EISCAT_3D concept



HPC-approach needed!

Data available:

- "cp6" — from 1990-06-13 to 1998-09-24, 76 experiments, ~52 full days (full day = 24h)
- "arcd" —from 2003-05-29, 183 experiments, ~110 full days
- "manda" — from 2005-04-14, more than 411 experiments, ~202 full days

Total duration of data available = **364 full days**

Full computational cost

- MCMC run time for each altitude = **~3.5 sec**
- No of altitude steps = **250**
- Computational time for each 5 min average profile = $3.5 \times 250 = \mathbf{875 \text{ sec} (\sim 15 \text{ min})}$
- 5 min-integrated profiles in 24 hr = **288**

Estimated computational time :

$364 \text{ full days} \times 288 \text{ 5-min-integrated profiles} \times 15 \text{ minutes per one profile} = 1572480 \text{ minutes} = \sim \mathbf{3 \text{ years of computational time!}}$

Incoherent scatter radar data levels

- Level 0: raw voltage samples
- Level 1: lag-profiles
- Level 2: plasma parameters (density, temperature, composition)
- Level 3: particle precipitation, conductivities, currents ...

Step 1: Lag-profiles to plasma parameters (level 1 to 2)

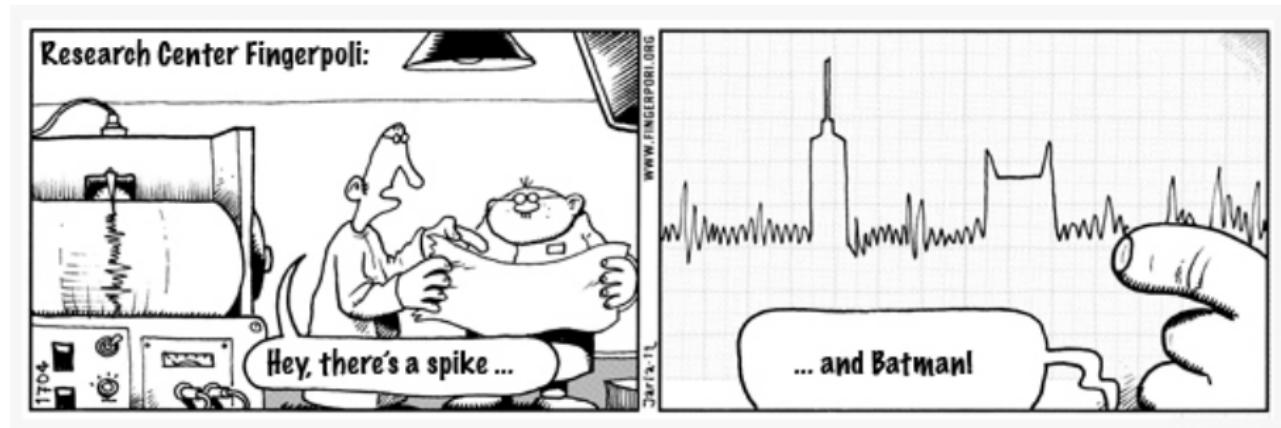


Figure: Fingerpori, HS, 7 September 2012

Incoherent scatter spectrum

The spectral density of the incoherent scattering is

$$\sigma(\omega_0 + \omega)d\omega = \frac{N_e r_e^2 d\omega}{\pi} \frac{\left(|y_e|^2 \frac{\sum_j n_j \Re(y_j)}{\omega - \mathbf{k} \cdot \mathbf{v}_{dj}} \left| \sum_j \mu_j y_j + i \lambda_D^2 k^2 \right|^2 \frac{\Re(y_e)}{\omega - \mathbf{k} \cdot \mathbf{v}_{de}} \right)}{\left(\left| y_e + \sum_j \mu_j y_j + i \lambda_D^2 k^2 \right|^2 \right)}.$$

Here

- $n_j = N_j / N_e$ and $\mu_j = n_j T_e / T_j$ (densities and temperatures),
- $\lambda_D = (\epsilon_0 k_B T_e / N_e e^2)^{1/2}$ (Debye length)
- $k = 2\pi/\lambda$ (wave number).

Incoherent scatter spectrum: Inputs

What is needed for calculating the IS spectrum?

- Radio wave parameters: ω , \mathbf{k}
- Plasma composition, temperatures, ion masses, coll. frequencies and drift velocities: N_j , T_j , m_j , ν_j and \mathbf{v}_j
- Magnetic field \mathbf{B}

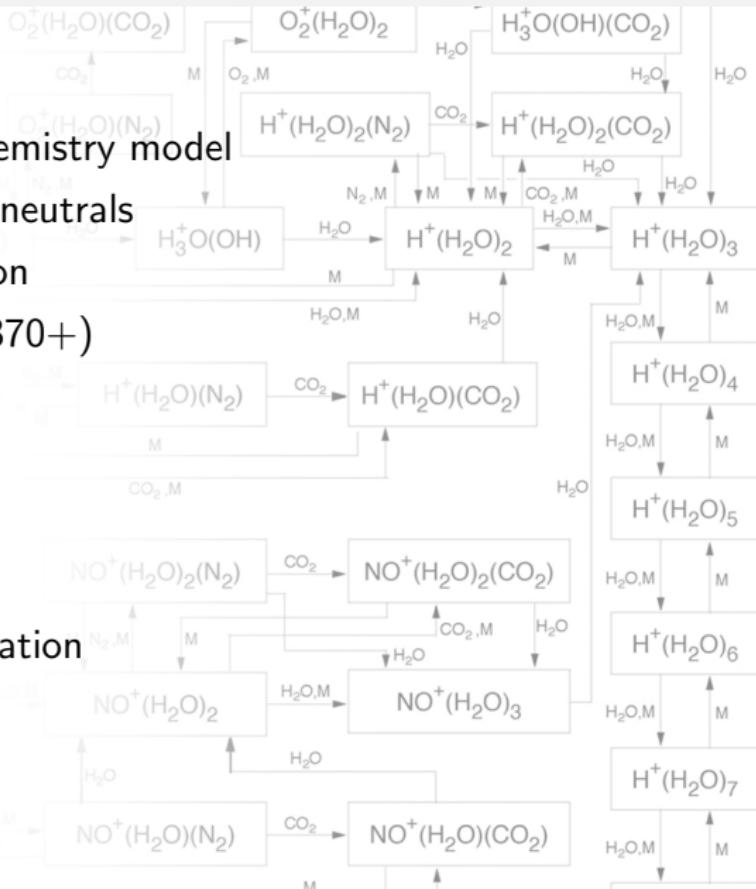
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Detailed 1-D time dependent chemistry model

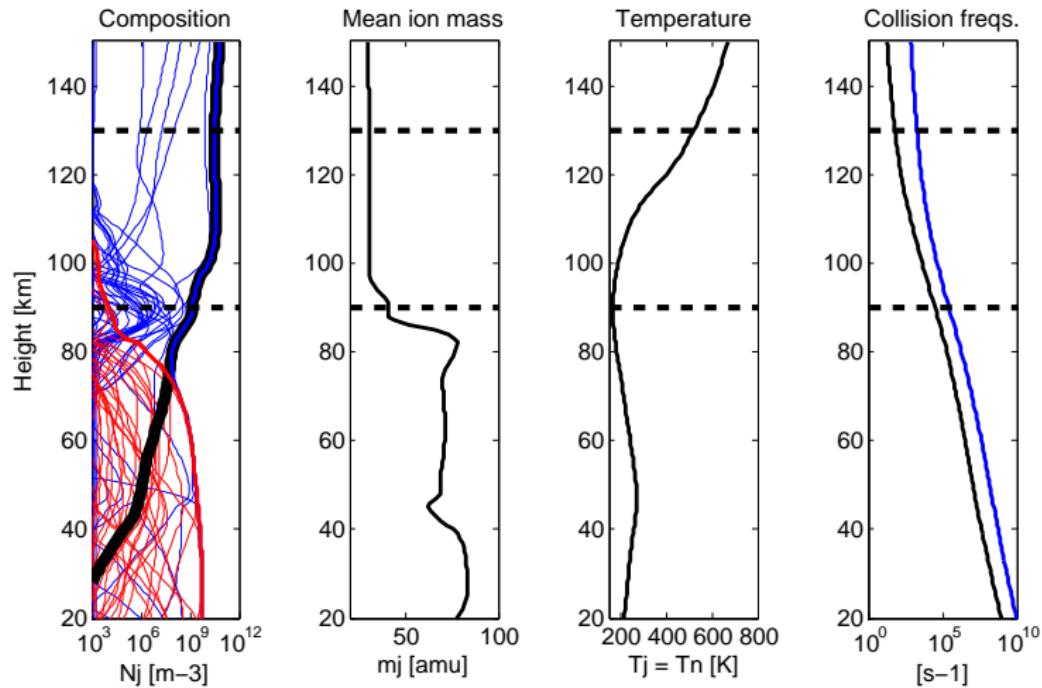
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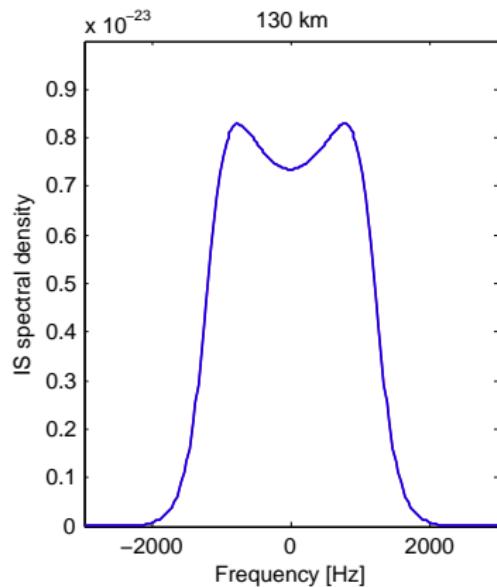
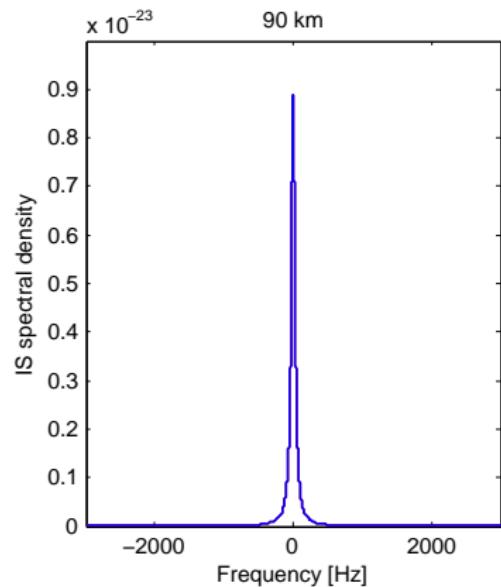
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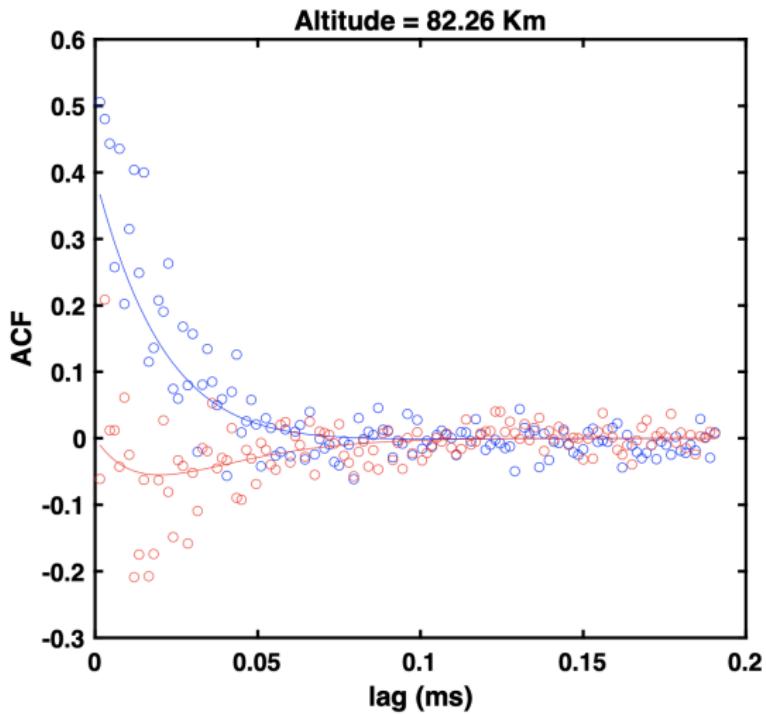
Input profiles (based on the SIC model)



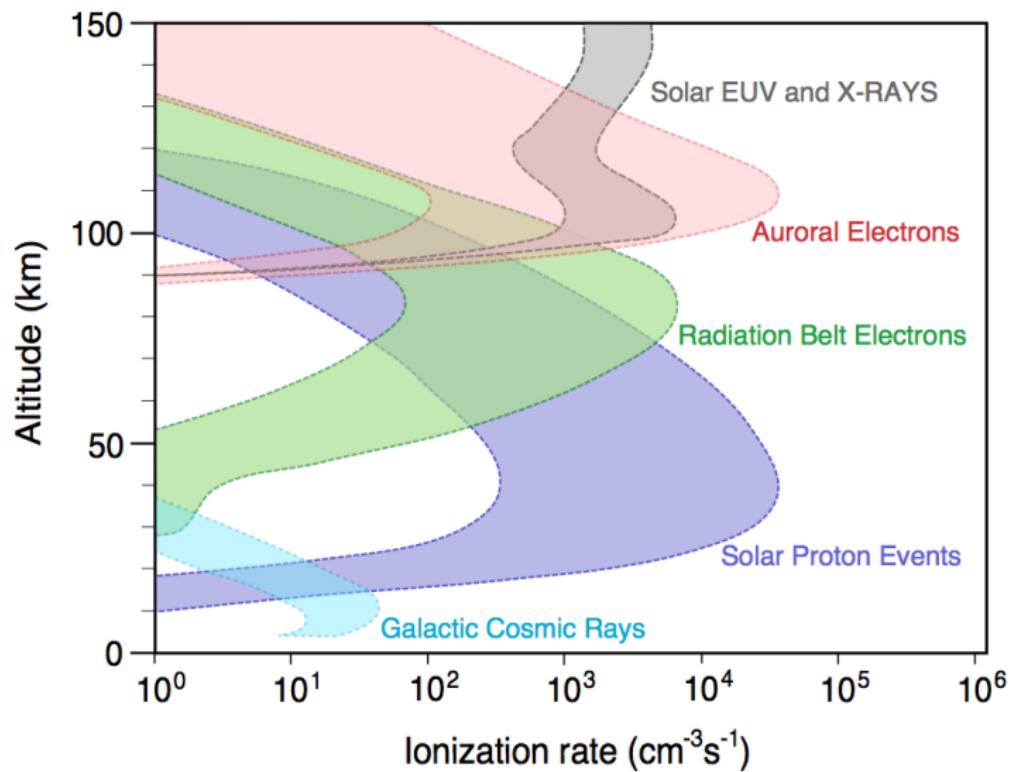
Incoherent scatter ion line: 90 and 130 km



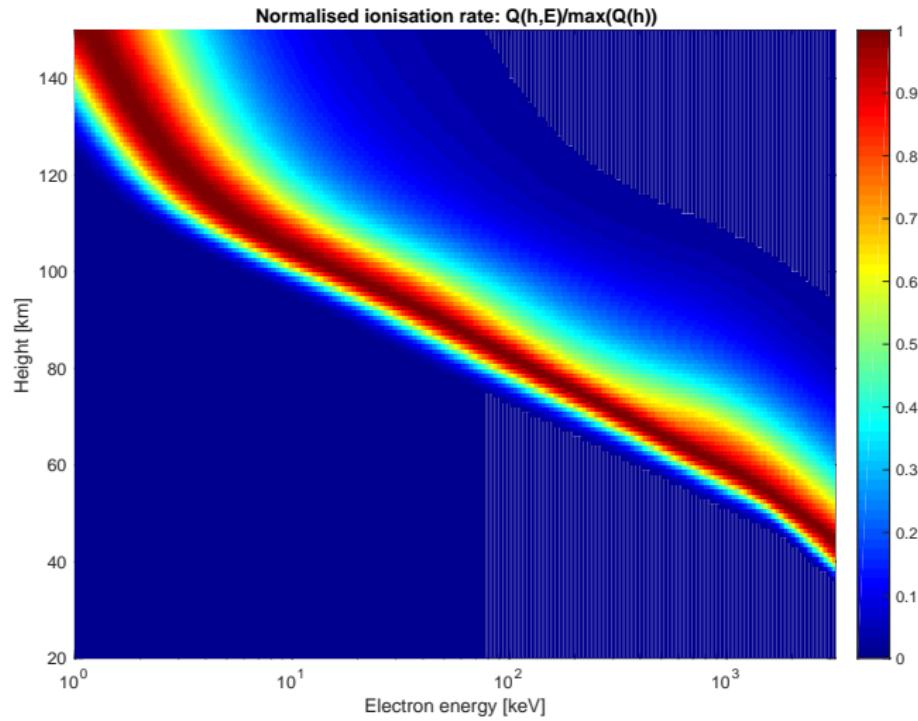
Plasma parameter fit to the ACFs measured



Step 2: Energetic particle precipitation inversion

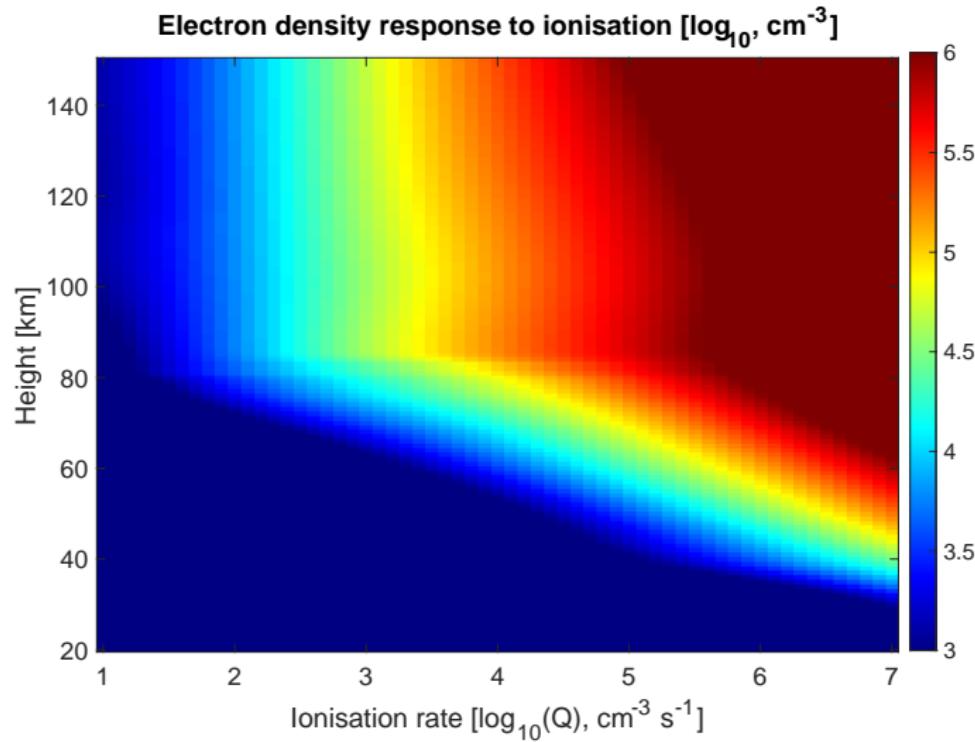


Electron precipitation driven ionisation: height vs. altitude



¹Fang et al., Parameterization of monoenergetic electron impact, GRL, 2010

Ionisation impact on electron density (SIC model)



EPP inversion:

$$\phi(E) \rightarrow Q(h) \rightarrow N_e(h)$$

① $\phi(E) \rightarrow Q(h)$

- ▶ Calculate $Q(h) = \int A(h, E) \phi(E) dE$
- ▶ Matrix $A(h, E)$ contains the atmospheric response²

② $Q(h) \rightarrow \text{SIC} \rightarrow N_e(h)$

- ▶ Pre-calculate N_e at a sufficiently fine grid of Q and interpolate $Q \rightarrow N_e$

③ Inversion

- ▶ Proposal $\phi(E)$ is an arbitrary piece-wise log-log (here: 7 nodes)
- ▶ Minimise in MCMC chain ($N=100\ 000$):
$$\chi^2 = \sum_h ((\log(N_e(h)) - \log(N_{e:\text{data}}(h))) / W)^2$$

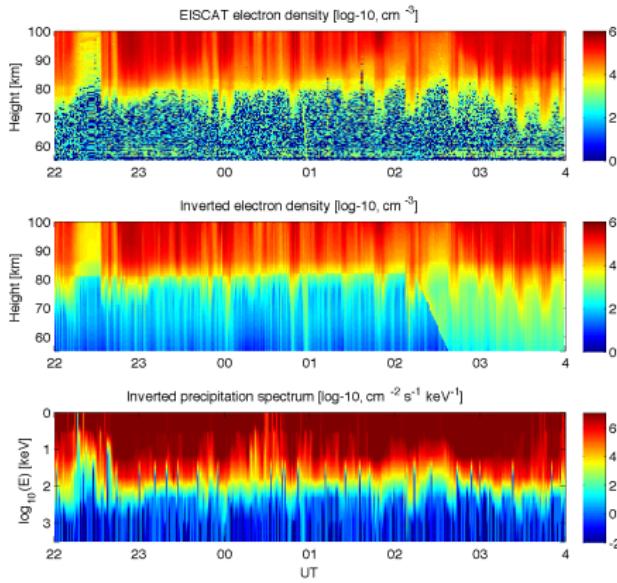
¹Fang, X., C. E. Randall, D. Lummerzheim, W. Wang, G. Lu, S. C. Solomon, and R. A. Frahm (2010), Parameterization of monoenergetic electron impact ionization, Geophys. Res. Lett., 37, L22106, doi:10.1029/2010GL045406.

Why MCMC inversion?

- A non-linear problem
- Easy to construct the χ^2 error function (e.g., in log-scale)
- Easy to limit the solution (e.g., only positive fluxes are allowed)
- Kick-ass MCMC toolbox available¹

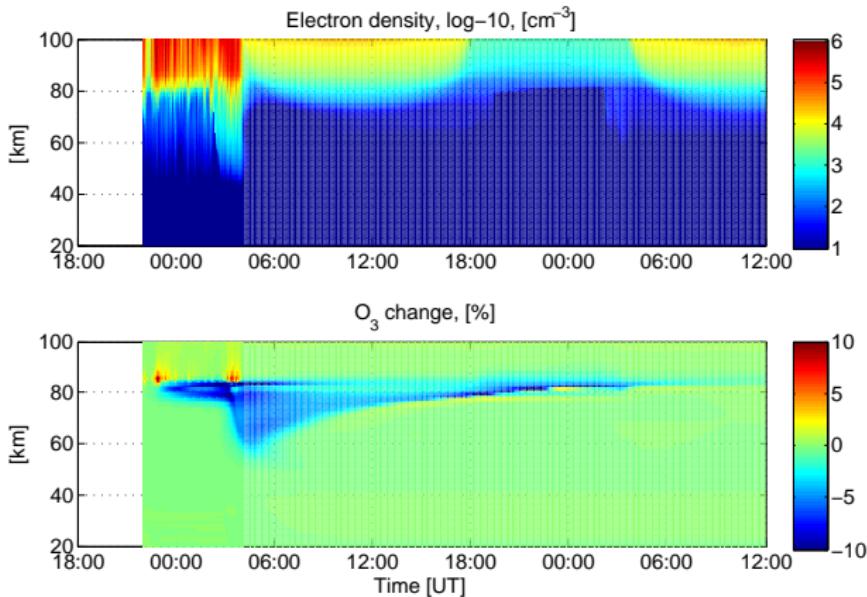
¹Search: 'Marko Laine + MCMC' → <http://helios.fmi.fi/~lainema/mcmc/>

Example case: 28–29 March 2017



¹Miyoshi et al. (2021), Penetration of MeV electrons into the mesosphere accompanying pulsating aurorae, *Scientific Reports* (Springer-Nature).
<https://doi.org/10.1038/s41598-021-92611-3>

Example case: 28–29 March 2017



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Summary of part 3/3

- Inversion of the precipitation characteristics (ϕ, Q, N_e) from the incoherent scatter radar data
- Full time-dependent ion chemistry of the SIC model is used in the forward model → chemical consequences
- Results provide means to constrain better the particle forcing input to the climate models (such as WACCM-D)